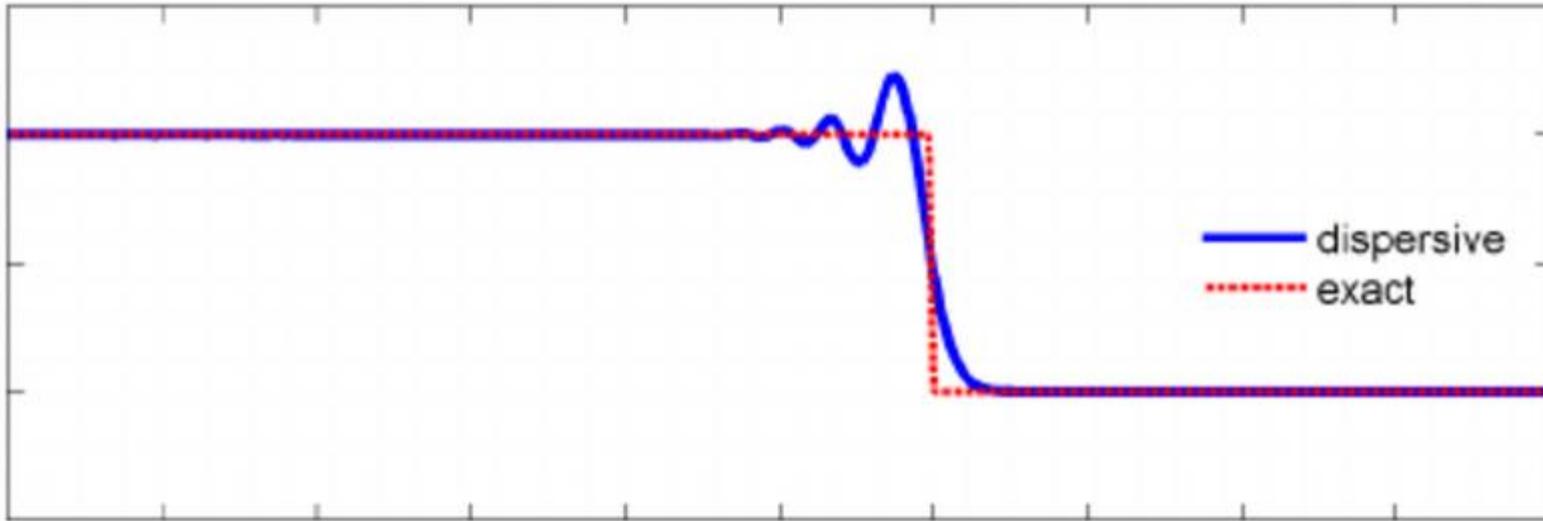


SCHISM physical formulation

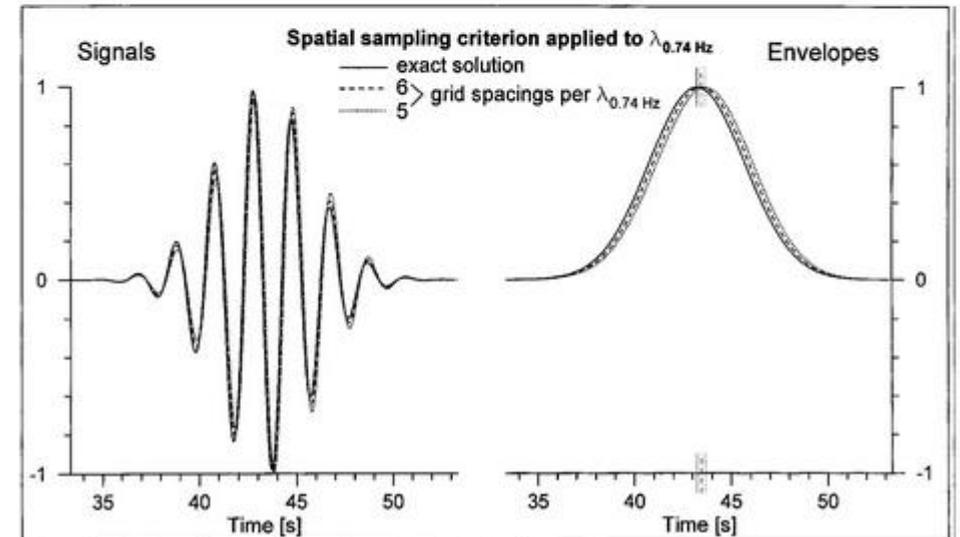
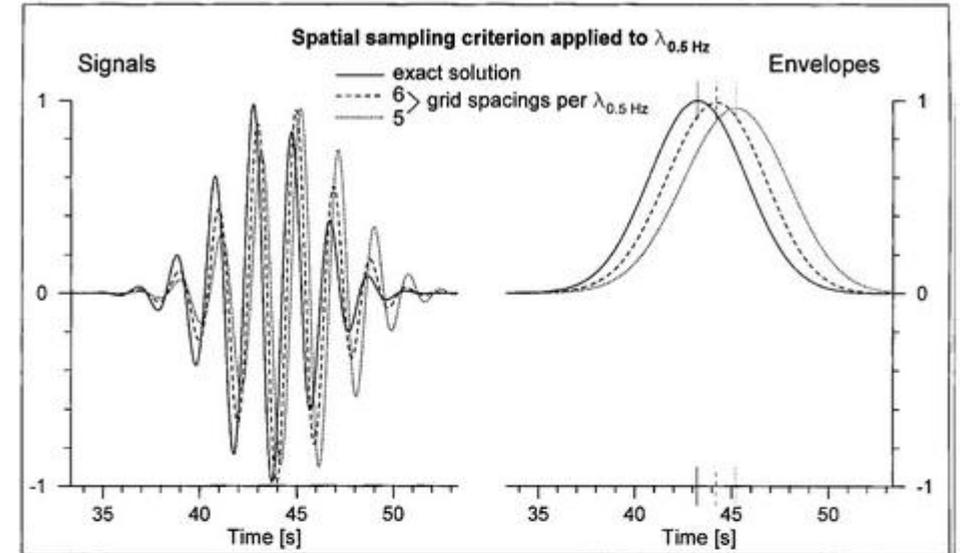
Joseph Zhang

Dispersion



- Waves of different frequencies travel at different phase speeds
- Usually related to derivatives of odd order (of either physical or numerical origin)

$$\varphi_t + \varphi_{xxx} - 6\varphi\varphi_x = 0$$



With dispersion

No dispersion

Governing equations: Reynolds-averaged Navier-Stokes (with vegetation)

- Continuity equation

$$\nabla \cdot \mathbf{u} + \frac{\partial w}{\partial z} = 0, \quad (\mathbf{u} = (u, v))$$

- Momentum equations

$$\frac{D\mathbf{u}}{dt} = \mathbf{f} - g\nabla\eta + \mathbf{m}_z - \alpha|\mathbf{u}|\mathbf{u}L(x, y, z)$$

vegetation

$$\mathbf{f} = f(v, -u) - \frac{g}{\rho_0} \int_z^\eta \nabla \rho d\zeta - \frac{\nabla p_A}{\rho_0} + \alpha g \nabla \Psi + \mathbf{F}_m + \text{other}$$

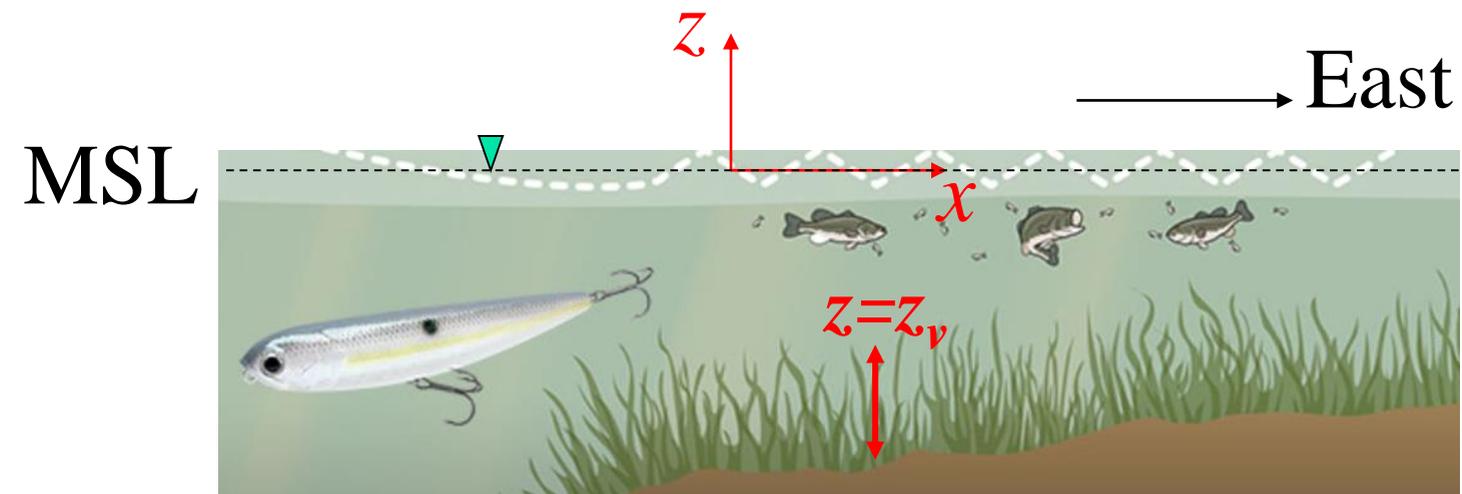
Vertical b.c. (3D):

$$\begin{cases} v \frac{\partial \mathbf{u}}{\partial z} = \boldsymbol{\tau}_w, & z = \eta \\ v \frac{\partial \mathbf{u}}{\partial z} = \chi \mathbf{u}_b, & z = -h \end{cases}$$

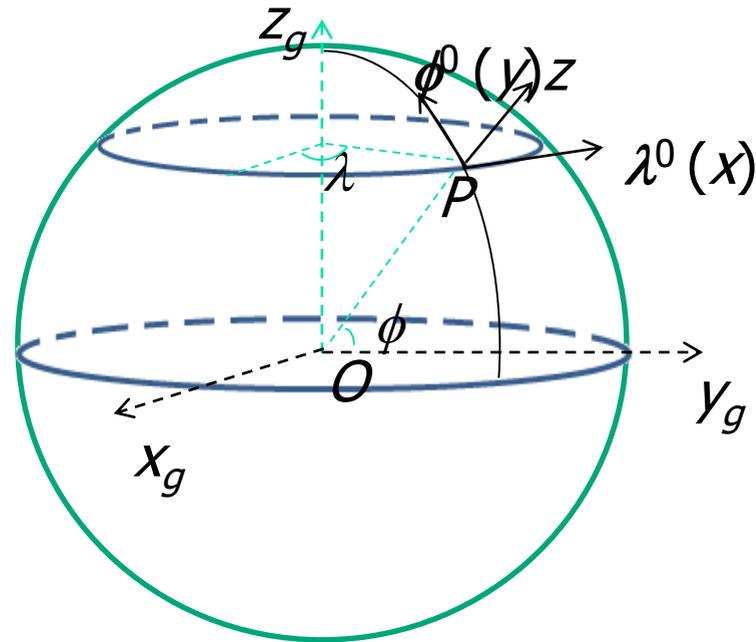
$$\mathbf{m}_z = \begin{cases} \frac{\partial}{\partial z} \left(v \frac{\partial \mathbf{u}}{\partial z} \right), & 3D \\ \frac{\boldsymbol{\tau}_w - \chi \mathbf{u}}{H}, & 2D \end{cases}$$

$$L(x, y, z) = \begin{cases} \mathcal{H}(z_v - z), & 3D \\ 1, & 2D \end{cases}$$

$$\alpha(x, y) = D_v N_v C_{Dv} / 2 \quad (\text{vegetation})$$



Spherical coordinates



We transform the coordinates instead of eqs

There are 2 frames used:

1. Global
2. Lon/lat (local @ node/element/side)

$$\begin{cases} x_g = R_0 \cos \phi \cos \lambda \\ y_g = R_0 \cos \phi \sin \lambda \\ z_g = R_0 \sin \phi \end{cases}$$

$$\lambda^0 = -\sin \lambda \mathbf{i} + \cos \lambda \mathbf{j}$$

$$\phi^0 = -\cos \lambda \sin \phi \mathbf{i} - \sin \lambda \sin \phi \mathbf{j} + \cos \phi \mathbf{k}$$

- Avoids polar singularity easily
- Preserves all original matrix properties

Changes in the code is mainly related to re-project vectors and to calculate distances

Hydrostatic model

Hydrostatic assumption

$$-\frac{1}{\rho} \frac{\partial P}{\partial z} - g = 0 \Rightarrow P = g \int_z^\eta \rho d\zeta + P_A$$

Separation of horizontal and vertical scales

$$\frac{D\mathbf{u}}{Dt} = -\frac{1}{\rho} \nabla \left(\int_z^\eta \rho d\zeta + P_A \right) + \nabla \cdot (\mu \nabla \mathbf{u}) + \frac{\partial}{\partial z} \left(\nu \frac{\partial \mathbf{u}}{\partial z} \right) - f \mathbf{k} \times \mathbf{u}$$

$$\frac{D\mathbf{u}}{Dt} = -g \nabla \eta + \frac{\partial}{\partial z} \left(\nu \frac{\partial \mathbf{u}}{\partial z} \right) - f \mathbf{k} \times \mathbf{u} + \alpha g \nabla \psi - \frac{1}{\rho_0} \nabla P_A - \frac{g}{\rho_0} \int_z^\eta \nabla \rho d\zeta + \nabla \square (\mu \nabla \mathbf{u})$$

Boussinesq assumption (\Rightarrow incompressibility)

$$\nabla \int_z^\eta \rho d\zeta = \int_z^\eta \nabla \rho d\zeta + \rho \nabla \eta$$

Momentum equation: vertical boundary condition (b.c.)

Surface

$$\nu \frac{\partial \mathbf{u}}{\partial z} = \boldsymbol{\tau}_w \quad \text{at } z = \eta$$

Bottom

$$\nu \frac{\partial \mathbf{u}}{\partial z} = C_D |\mathbf{u}_b| \mathbf{u}_b, \quad \text{at } z = -h$$

- Logarithmic law:

$$\mathbf{u} = \frac{\ln[(z+h)/z_0]}{\ln(\delta_b/z_0)} \mathbf{u}_b, \quad (z_0 - h \leq z \leq \delta_b - h)$$

z_0 : bottom roughness

- Reynolds stress:

$$\nu \frac{\partial \mathbf{u}}{\partial z} = \frac{\nu}{(z+h) \ln(\delta_b/z_0)} \mathbf{u}_b$$

$$\nu = \sqrt{2} s_m K^{1/2} l,$$

$$s_m = g_2,$$

$$K = \frac{1}{2} B_1^{2/3} C_D |\mathbf{u}_b|^2$$

$$l = \kappa_0 (z+h)$$

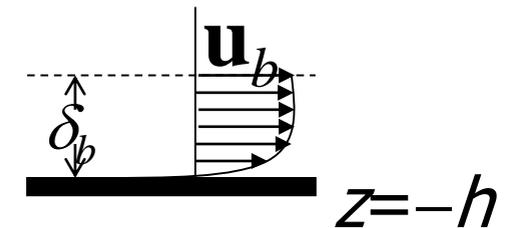
- Turbulence closure:

- Reynolds stress (const.)

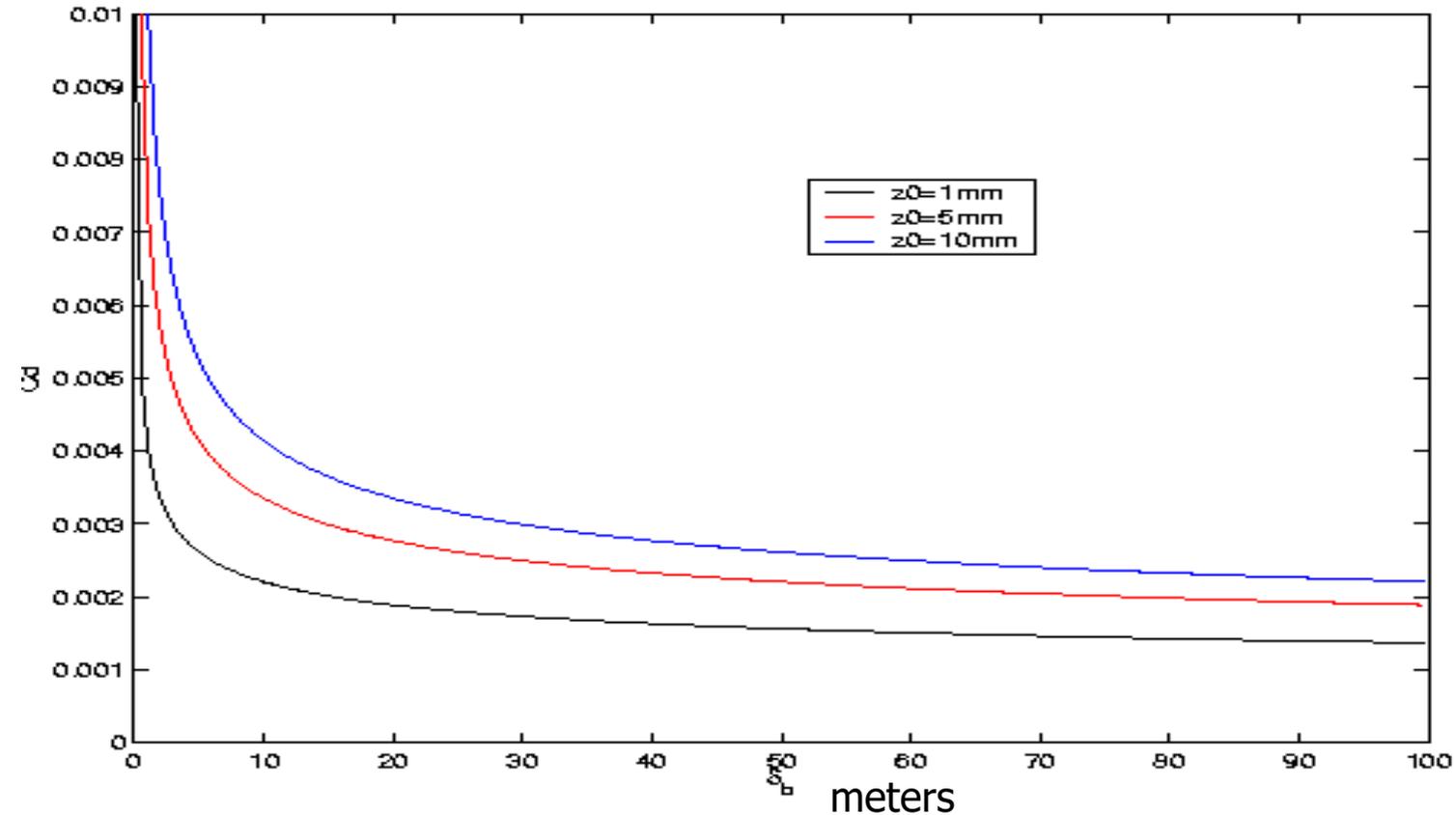
$$\nu \frac{\partial \mathbf{u}}{\partial z} = \frac{\kappa_0}{\ln(\delta_b/z_0)} C_D^{1/2} |\mathbf{u}_b| \mathbf{u}_b, \quad (z_0 - h \leq z \leq \delta_b - h)$$

- Drag coefficient:

$$C_D = \left(\frac{1}{\kappa_0} \ln \frac{\delta_b}{z_0} \right)^{-2}$$



Bottom drag



- Use of z_0 seems most natural option, but tends to over-estimate C_D in shallow area
- C_D should vary with bottom layer thickness (i.e. vertical grid)
- Different from 2D, 3D results of elevation depend on vertical grid (with constant C_D)
- Simplest option for compound flooding study is to use variable Manning's n (2D)

Wind Shear Stress: Turbulent Flux of Momentum

Calculation of shear stress follows naturally from calculation of turbulent heat fluxes

Total shear stress of atmosphere upon surface:

$$\tau_w = \frac{\rho_a}{\rho_0} u_*^2$$

Alternatively, Pond and Picard's formulation can be used as a simpler option

$$\boldsymbol{\tau}_w = \frac{\rho_a}{\rho_0} C_{ds} |\mathbf{u}_w| \mathbf{u}_w$$

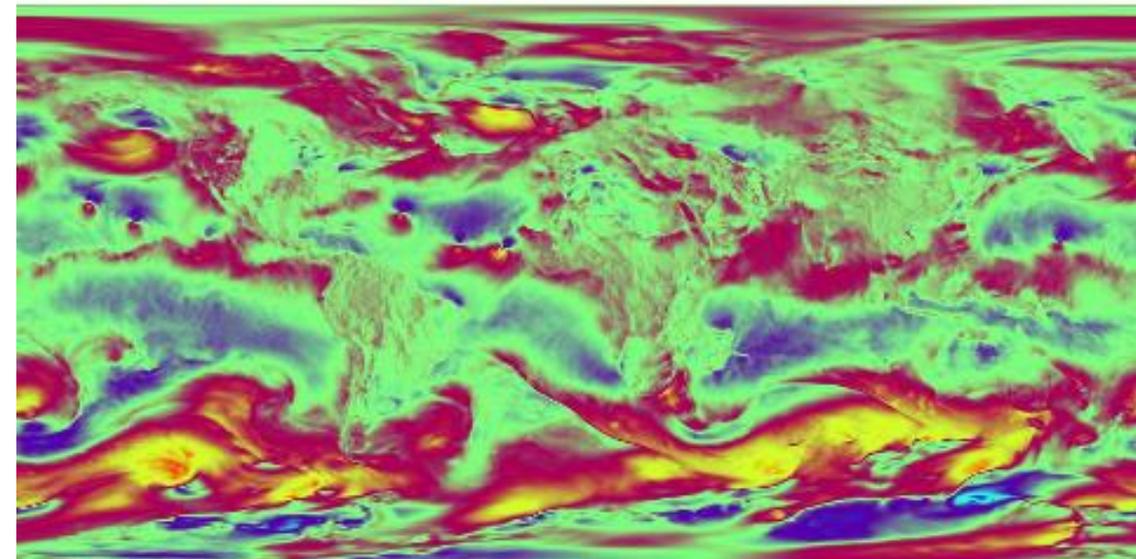
$$C_{ds} = \frac{0.61 + 0.063 u_w'}{1000}$$

$$u_w' = \max(6, \min(50, u_w))$$

Atmospheric inputs to air-sea exchange module

- Forecasts of u_w v_w @ 10m, P_A @ MSL, T_{air} and q_{air} @2m
 - used by the bulk aerodynamic model to calculate the scaling parameters
 - T_{air} and q_{air} not used in 2D; so set a constant for these
- Precipitation (sflux_prc*.nc): must be converted to source/sink for 2D config
- We have scripts to convert several products into our sflux inputs (CF1.0 convention; structured grids)
 - CFSR
 - ERA
 - GFS
- Atmospheric forcing is also passed onto wave module (WWM) when invoked

Global ERA5 (u-wind)



Classification of six types of culvert flow

TYPE	EXAMPLE	TYPE	EXAMPLE
<p>1 CRITICAL DEPTH AT INLET</p> <p>$\frac{h_1 - z}{D} < 1.5$ $h_4/h_c < 1.0$ $S_0 > S_c$</p>	<p>$Q = CA_c \sqrt{2g(h_1 - z + a_1 \frac{V_1^2}{2g} - d_c - h_{f_{1,2}})}$</p>	<p>4 SUBMERGED OUTLET</p> <p>$\frac{h_1 - z}{D} > 1.0$ $h_4/D > 1.0$</p>	<p>$Q = CA_0 \sqrt{\frac{2g(h_1 - h_4)}{1 + \frac{29C^2n^2L}{R_0^{4/3}}}}$</p>
<p>2 CRITICAL DEPTH AT OUTLET</p> <p>$\frac{h_1 - z}{D} < 1.5$ $h_4/h_c < 1.0$ $S_0 < S_c$</p>	<p>$Q = CA_c \sqrt{2g(h_1 + a_1 \frac{V_1^2}{2g} - d_c - h_{f_{1,2}} - h_{f_{2,3}})}$</p>	<p>5 RAPID FLOW AT INLET</p> <p>$\frac{h_1 - z}{D} \approx 1.5$ $h_4/D \approx 1.0$</p>	<p>$Q = CA_0 \sqrt{2g(h_1 - z)}$</p>
<p>3 TRANQUIL FLOW THROUGHOUT</p> <p>$\frac{h_1 - z}{D} < 1.5$ $h_4/D \leq 1.0$ $h_4/h_c > 1.0$</p>	<p>$Q = CA_3 \sqrt{2g(h_1 + a_1 \frac{V_1^2}{2g} - h_3 - h_{f_{1,2}} - h_{f_{2,3}})}$</p>	<p>6 FULL FLOW FREE OUTFALL</p> <p>$\frac{h_1 - z}{D} \approx 1.5$ $h_4/D \approx 1.0$</p>	<p>$Q = CA_0 \sqrt{2g(h_1 - h_3 - h_{f_{2,3}})}$</p>

There is a manual for hydraulics structures in SCHISM

Orifice equation

$$Q = CA \sqrt{2g(y_1 - y_3)}$$

where $A =$ Area of the gate opening
 $C =$ Discharge coefficient (typically 0.8)